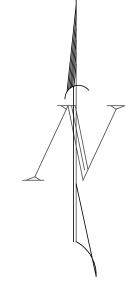
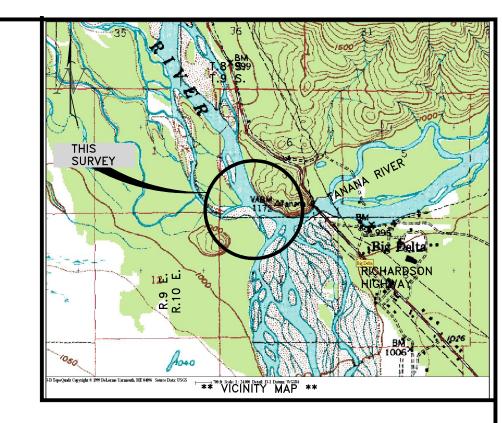
# **EXHIBIT G**

# PROJECT BOUNDARY MAPS

The project boundary is within that granted under the preliminary permit issued to WPC under Project No. 13305 and is shown below. The exact location of the device within the project boundary is proposed to be 64°09'22.66" N, 145°51'39.88" W on the right bank of the Tanana River near the community of Whitestone.







## \*\* SURVEY NOTES \*\*

- 1. THE BASIS OF BEARINGS FOR THIS SURVEY WAS THE COMPUTED LINE BETWEEN N.G.S. STATION "PROBERT" LAT.64°08'50.93678"N LONG.145°49'54.80231"W ALASKA STATE PLANE COORDINATES N3708250.4987, E1667259.6450 AND A RTK/GPS BASE STATION LAT.64°08'25.85756"N LONG. 145°50'21.49237"W ALASKA STATE PLANE COORDINATES N3705699.5358, E1666082.1594 THE COMPUTED BEARING BEING S24°46'38"W.
- 2. COORDINATES REFERRED TO HEREON ARE EXPRESSED IN NAD 83 ALASKA STATE PLANE, ZONE 5003 CONVERTED TO HORIZONTAL GROUND US SURVEY FEET.
- 3. THE BASIS OF ELEVATION FOR THIS SURVEY WAS A RECOVERED BRASS DISK SET IN THE TOP OF A CONCRETE CURB, AT THE S.W. END OF THE RICHARDSON HIGHWAY BRIDGE HAVING A PURPORTED ELEVATION OF 998.94 FEET. DATUM IS UNKNOWN.
- 4. THIS SURVEY WAS ACCOMPLISHED USING RTK/GPS METHODS. EQUIPMENT CONSISTED OF SPECTRA EPOCH L1/L2 SURVEY GRADE RECEIVERS HORIZONTALLY/VERTICALLY CORRECTED TO THE BASIS OF BEARING/ELEVATIONS WITH TDS SURVEY PRO-GPS SOFTWARE.
- 5. THE HYDROGRAPHIC SURVEY WAS ACCOMPLISHED USING RTK/GPS METHODS INTERFACED WITH AN OHMEX ULTRA—HIGH FREQUENCY, NARROW BEAM, SURVEY GRADE SONAR SOUNDER CORRECTED IN REAL-TIME TO THE BASIS OF ELEVATIONS WITH TDS SURVEY PRO-GPS SOFTWARE AND PROCESSED IN THE OFFICE WITH HY-PAC SOFTWARE.

# \*\* PROPOSED PROJECT BOUNDARY COORDINATES \*\*

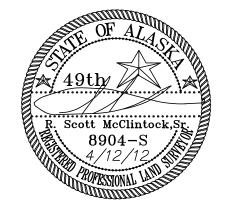
	APPROX. 23 ACRES	
	NORTH COORDINATE	EAST COORDINATE
A B C C	3712043.8380 3711426.0621 3711182.6491 3711231.7259 3711360.8461	1662206.7916 1663006.9732 1662812.8590 1661458.7297 1661089.5265
=	3711569.3368	1661162.8888

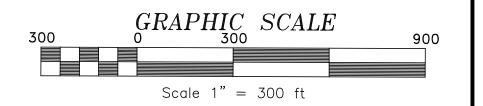
\* SURVEYOR'S CERTIFICATE \*

I HEREBY CERTIFY THAT I AM PROPERLY REGISTERED AND LICENSED TO PRACTICE LAND SURVEYING IN THE STATE OF ALASKA, THAT THESE DRAWINGS REPRESENTS A SURVEY MADE BY ME OR UNDER MY DIRECT SUPERVISION, THAT THE FEATURES SHOWN HEREON ACTUALLY EXIST AS DESCRIBED,
AND THAT ALL DIMENSIONS, RELATIVE BEARINGS, ELEVATIONS AND OTHER DETAILS ARE
CORRECT TO THE BEST OF MY KNOWLEDGE.

DATE: 4/12/12

R. SCOTT McCLINTOCK, SR. P.L.S.







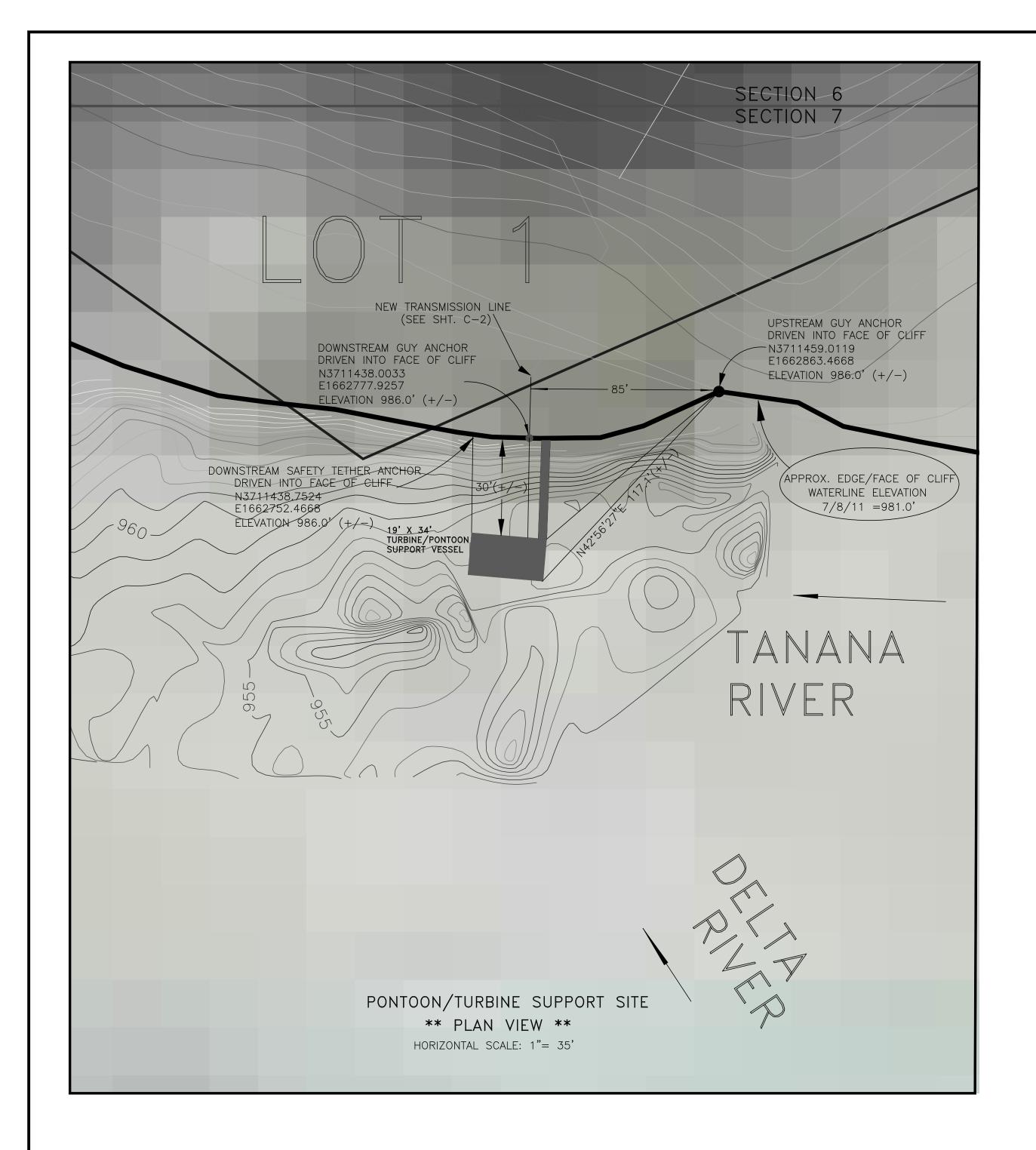
\* SURVEYING & MAPPING \*
P.O. BOX 1444 NOME, ALASKA 99762 (907) 443-6068 - SITE PLAN - TOPOGRAPHIC ELECTRIC ROUTE -- HYDROGRAPHIC SURVEY -

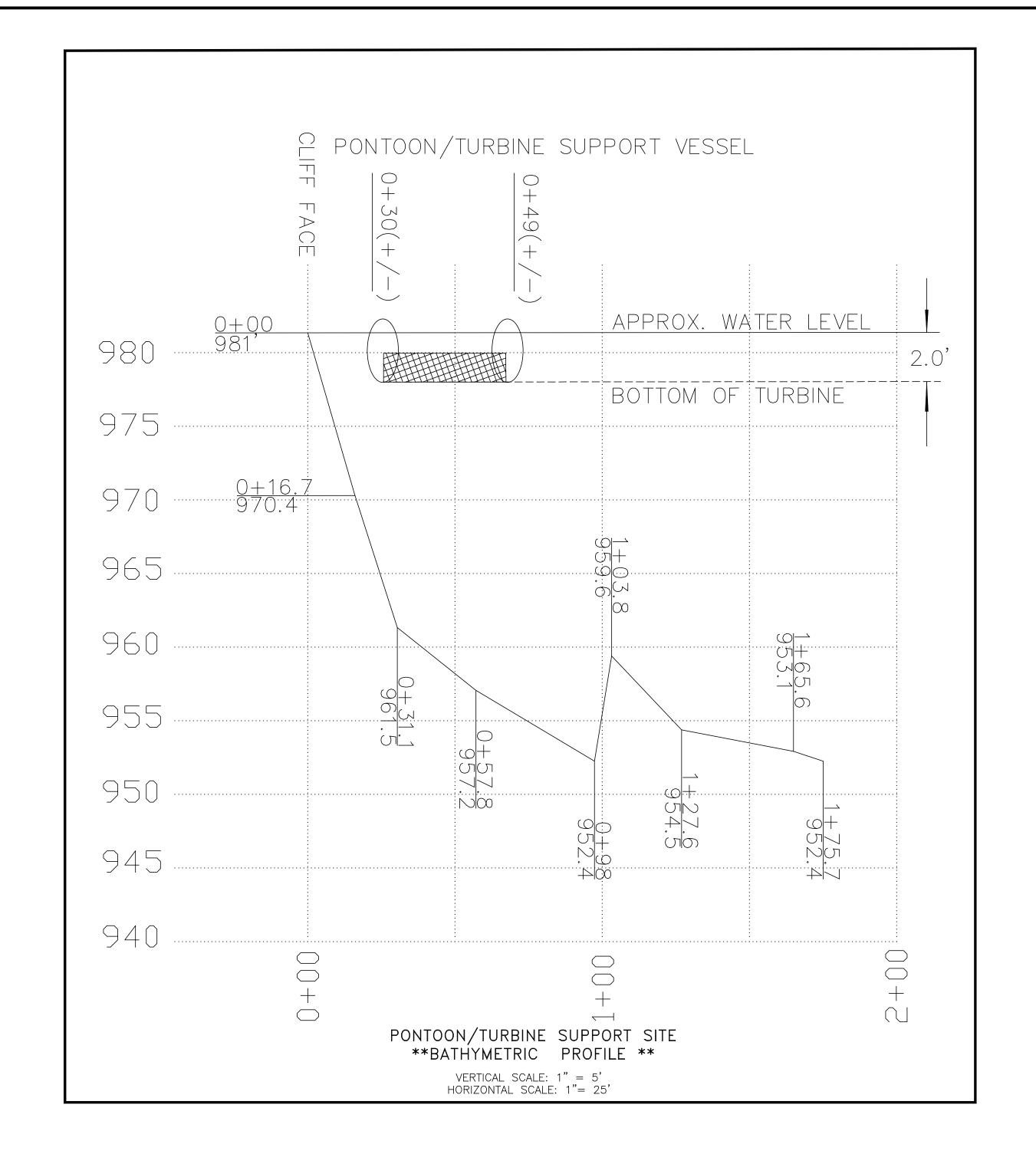
WHITESTONE POWER & COMMUNICATIONS PONCELET KINETICS RHK100 HYDROKINETIC TURBINE PROTOTYPE PROJECT within SECTIONS 6 & 7, T.9 S., R.10 E., FAIRBANKS MERIDIAN

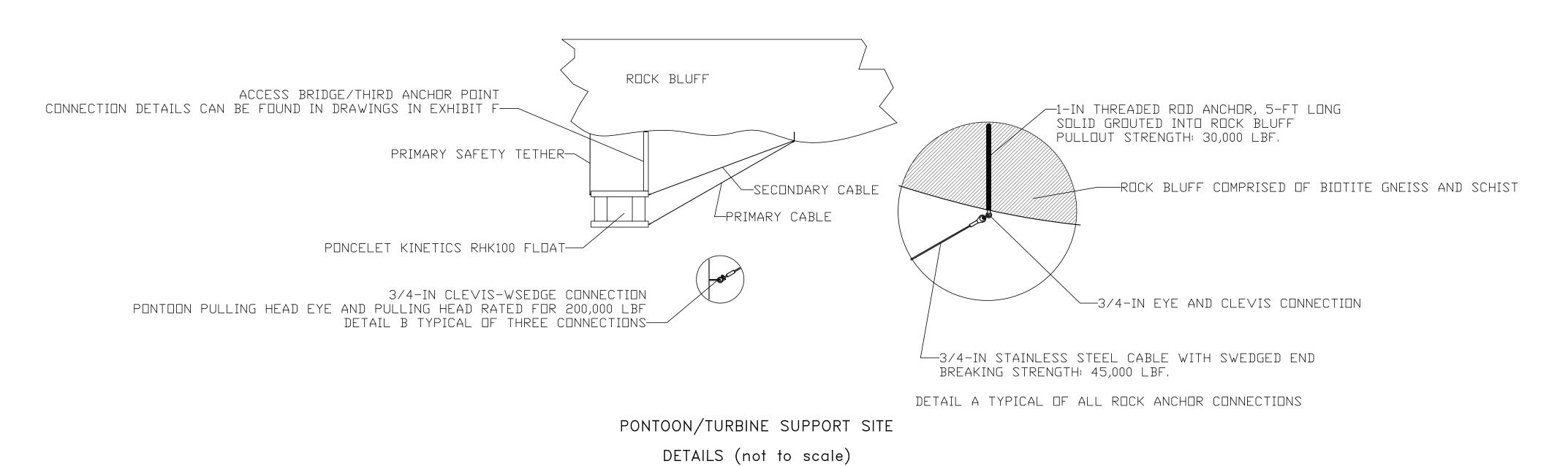
275.5 MILE, RICHARDSON HIGHWAY

BIG DELTA, ALASKA

WHITESTONE POWER & COMMUNICATIONS MR. STEVEN SELVAGGIO, REP. P.O. BOX 782 DELTA JUNCTION, ALASKA 99762 EXHIBIT G-2 AK11-011 D.N.R. FILE REF: XXXXXXXX-XXXX









\* SURVEYING & MAPPING \*
P.O. BOX 1444 NOME, ALASKA 99762 (907) 443-6068

- SITE PLAN - TOPOGRAPHIC ELECTRIC ROUTE - HYDROGRAPHIC SURVEY -

WHITESTONE POWER & COMMUNICATIONS

PONCELET KINETICS RHK100

HYDROKINETIC TURBINE PROTOTYPE PROJECT

within SECTIONS 6 & 7, T.9 S., R.10 E., FAIRBANKS MERIDIAN

275.5 MILE, RICHARDSON HIGHWAY

BIG DELTA, ALASKA

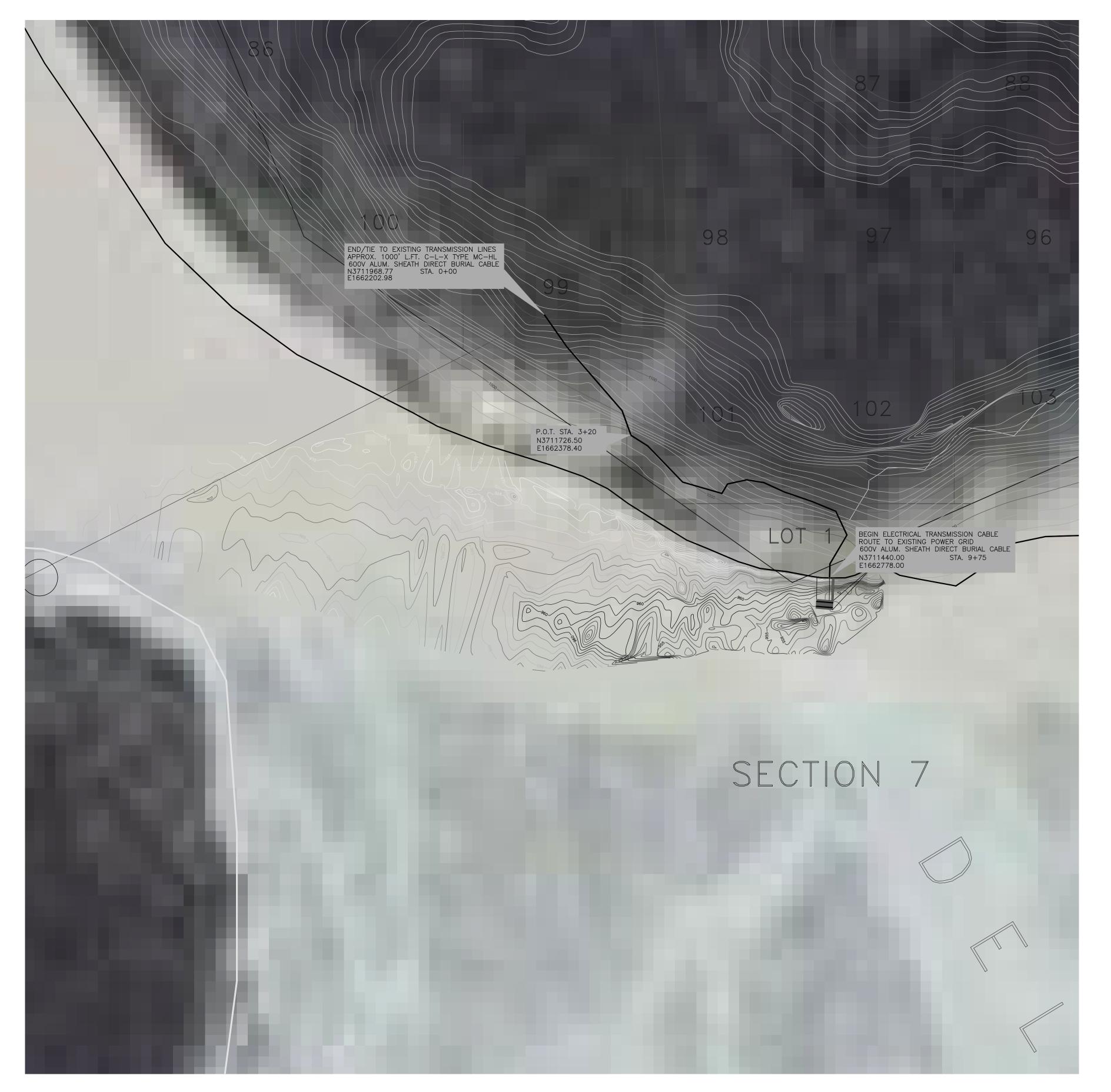
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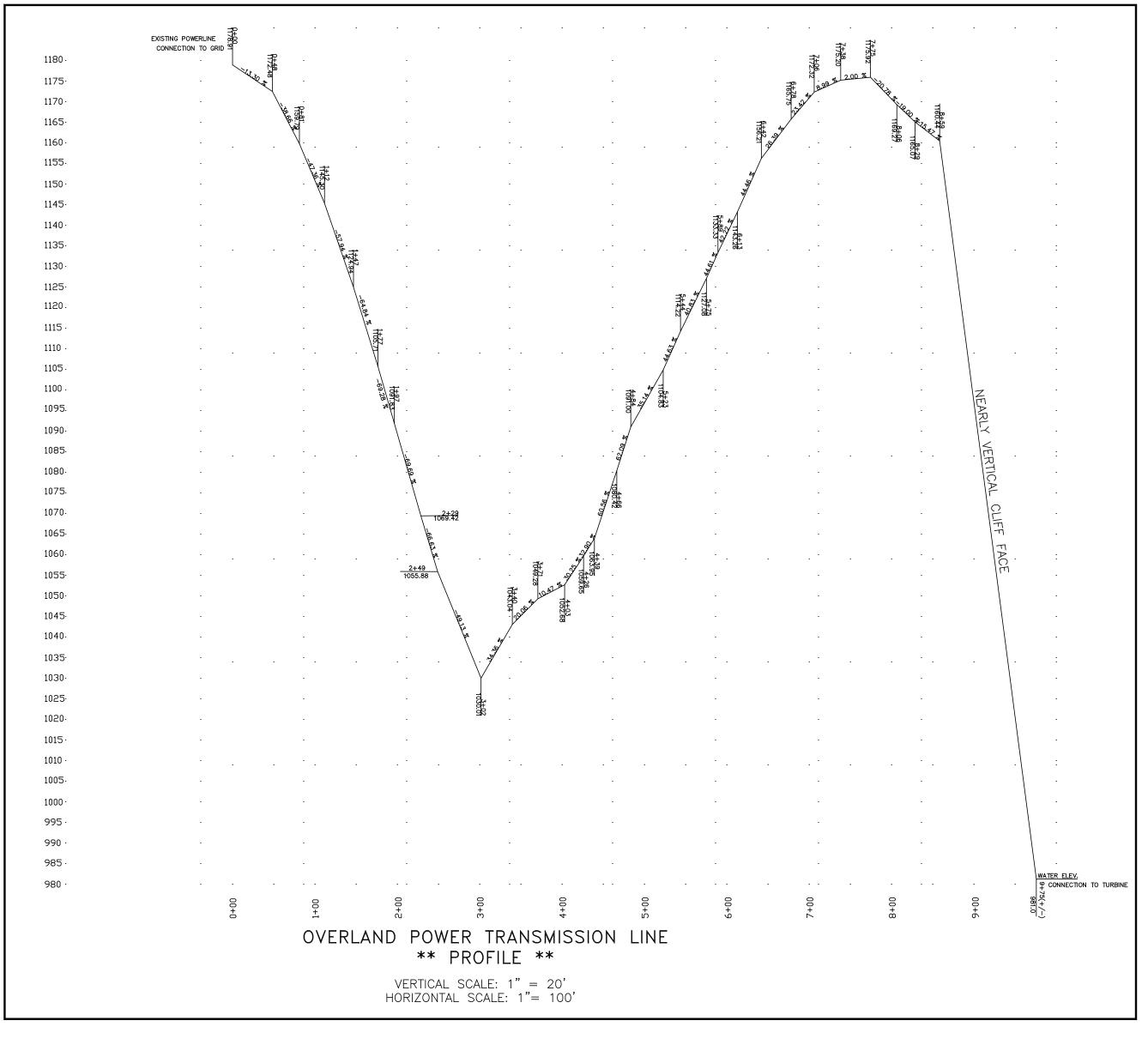
DATE:
APRIL 12, 2012

SHEET:
EXHIBIT G-3

DEVELOPER:
WHITESTONE POWER & COMMUNICATIONS
MR. STEVEN SELVAGGIO, REP.
P.O. BOX 782
DELTA JUNCTION, ALASKA 99762

D.N.R. FILE REF: XXXXXXXXX—XXXX





OVERLAND POWER TRANSMISSION LINE

\*\* PLAN VIEW \*\*

HORIZONTAL SCALE: 1"= 100'



\* SURVEYING & MAPPING \*
P.O. BOX 1444 NOME, ALASKA 99762 (907) 443-6068

- SITE PLAN - TOPOGRAPHIC ELECTRIC ROUTE - HYDROGRAPHIC SURVEY -

WHITESTONE POWER & COMMUNICATIONS

PONCELET KINETICS RHK100

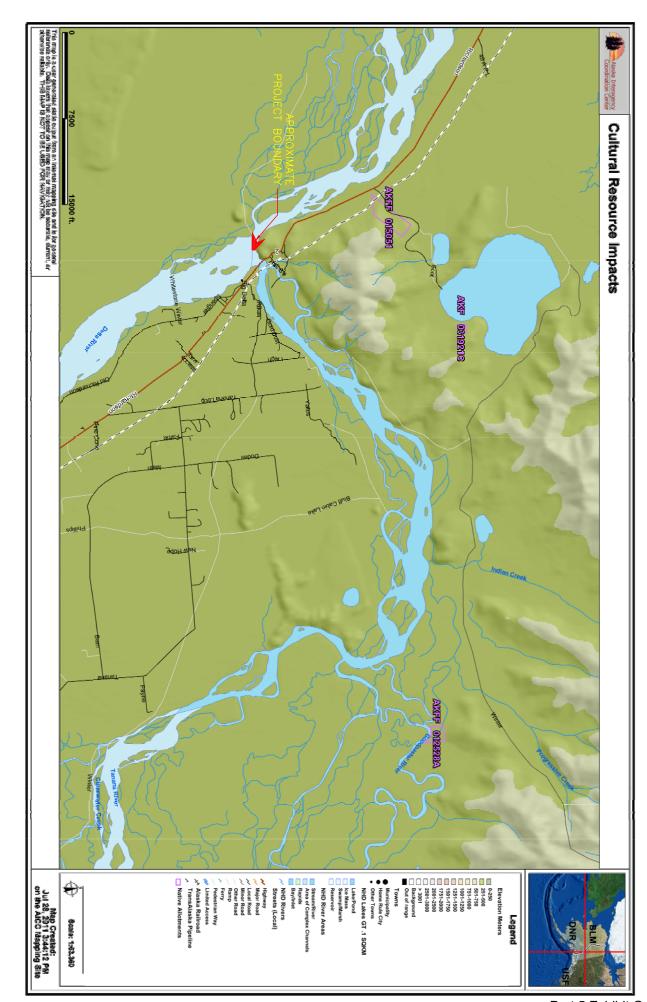
HYDROKINETIC TURBINE PROTOTYPE PROJECT

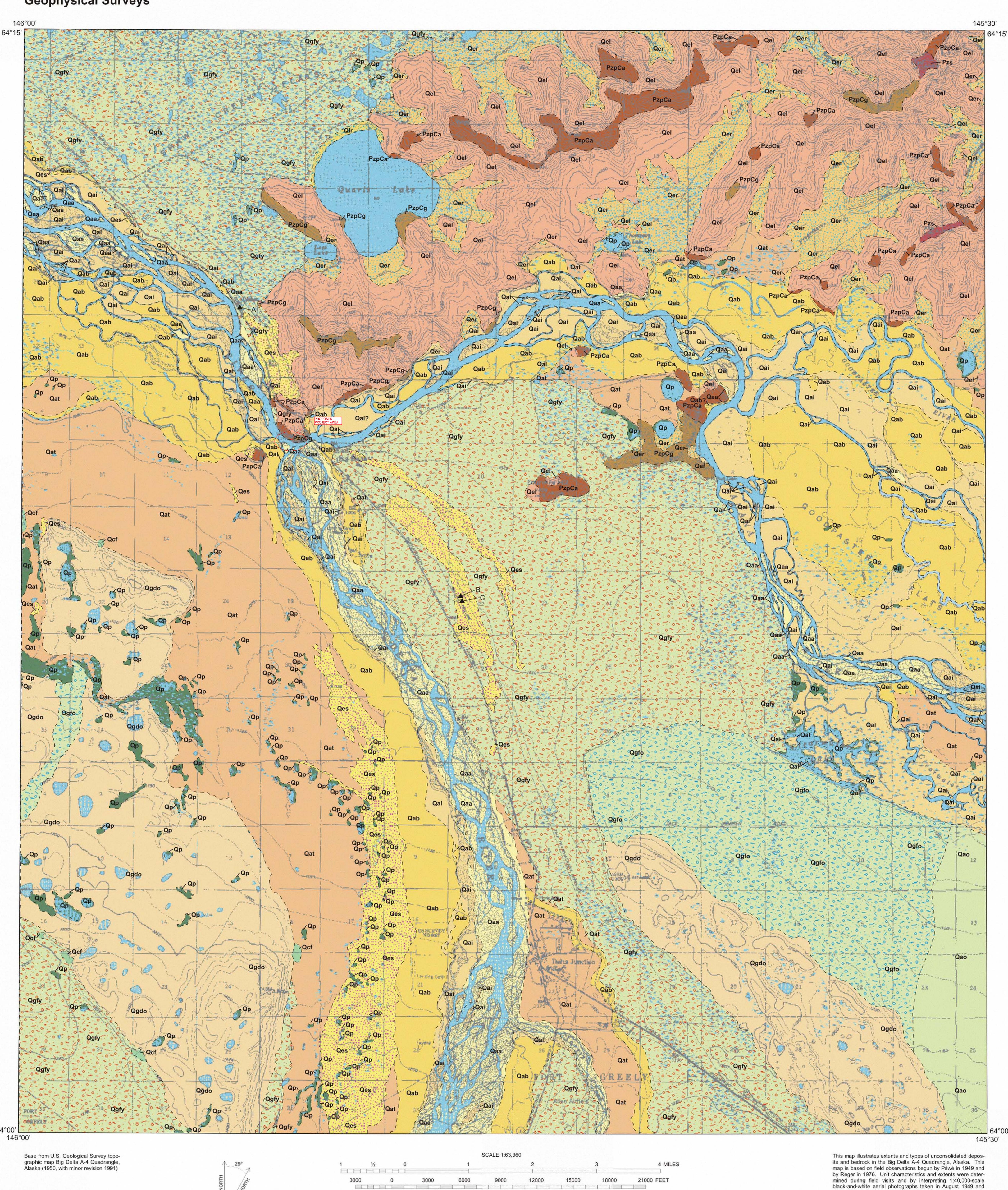
PONCELET KINETICS RHK100

HYDROKINETIC TURBINE PROTOTYPE PROJECT

within SECTIONS 6 & 7, T.9 S., R.10 E., FAIRBANKS MERIDIAN

275.5 MILE, RICHARDSON HIGHWAY





# DESCRIPTION OF MAP UNITS

# **Unconsolidated Deposits**

channels are floored by 5 to 20 ft of sand and silty sand that are generally unfrozen; fills of inactive channels include 7 to 12 ft of discontinuously frozen organic silt with moderate to high ice content over sand and grav-

ABANDONED-FLOODPLAIN ALLUVIUM—Chiefly 10 to 20 ft of overbank sandy silt and silty sand overlying sandy, polymictic riverbed gravel beneath surfaces with widespread lowland loess and local sand dunes and subject to stream flooding about once every 500 to 1,000 years (Mann and others, 1995); may include several surface levels; overbank sequences contain organic-silt channel fills 7 to 20 ft thick; surface peat generally

discontinuous to widespread; generally frozen with low to moderate ice content UNDIFFERENTIATED GLACIAL AND NONGLACIAL ALLUVIUM—Primarily coarse, polymictic gravel and sand deposited by former proglacial streams of the Delta glaciation and locally covered along courses of nonglacial streams by postglacial alluvium; may be terraced; locally mantled by loess and eolian sand; locally subject to

seasonal stream icings; sporadically frozen with low ice content

moisture content is low (3 to 4 percent); discontinuously frozen

STREAM-TERRACE ALLUVIUM—Chiefly 4 to more than 20 ft of organic sandy silt and silty sand overlying well-sorted, polymictic sand and gravel beneath stream terrace treads no longer subject to inundations by the stream that deposited the alluvium (Kreig and Reger, 1982); may include several levels and incorporate outwash alluvium in highest terraces; locally covered by up to 15 ft of lowland loess and eolian sand-blanket and dune complexes; locally subject to seasonal stream icings; continuously to discontinuously frozen with low to

KED COLLUVIUM AND ALLUVIAL DEPOSITS—Primarily fan-shaped, massive to poorly stratified fine material mixed with polymictic pebble and cobble gravel laid down by small, ephemeral streams draining glacial

lower south-facing slopes and continuously frozen and ice rich on some lower north-facing slopes and lowland

and eolian-cover deposits; discontinuously to continuously frozen with moderate to high ice content LOESS—Silt with up to 15 percent very fine sand carried by winds and deposited as a blanket over the downto floodplain sources; thickness ranges from >20 ft close to Delta River to ≈2 ft elsewhere (Lindholm and others, 1959); organic rich on lower slopes and lowland sites; moderate to high moisture content (more than 15 percent moisture) in lowland sites; generally unfrozen, except discontinuously frozen and ice rich on some

RETRANSPORTED LOESS AND LOWLAND SILT-Chiefly organic silt with variable amounts of sand and lenses of locally derived gravel that are deposited by slope runoff and seasonal streams draining bedrock debris-flow deposits; subject to seasonal stream and slope icings; discontinuously to continuously frozen with

DUNE SAND—Chiefly fine eolian sand with trace amounts (2 to 16 percent) of silt (Kreig and Reger, 1982, pl. 9); dunes stand 5 to 10 ft in relief; discontinuous with thicknesses up to ≈25 ft (Schoephorster, 1973); thicker and more extensive west of Delta River; generally covered by up to 2 ft of loess (Lindholm and others, 1959);

TILL AND ASSOCIATED MORAINAL DEPOSITS OF DELTA GLACIATION—Heterogeneous, nonstratified, polymictic pebble-cobble gravel with some sand and silt and few to numerous subangular to subrounded boulders deposited by glacial ice; discontinuously mantled by eolian sand and loess; discontinuously to continuously frozen with low to moderate ice content (Péwé and Holmes, 1964; Holmes, 1965)

OUTWASH ALLUVIUM—Polymictic sand and gravel deposited by former proglacial streams; may be terraced; mantled by loess and eolian sand; locally subject to seasonal stream icings; sporadically to discontinuously

OUTWASH OF DELTA GLACIATION (Péwé and Holmes, 1964; Holmes, 1965)

OUTWASH OF DONNELLY GLACIATION (Péwé and Holmes, 1964; Holmes, 1965)

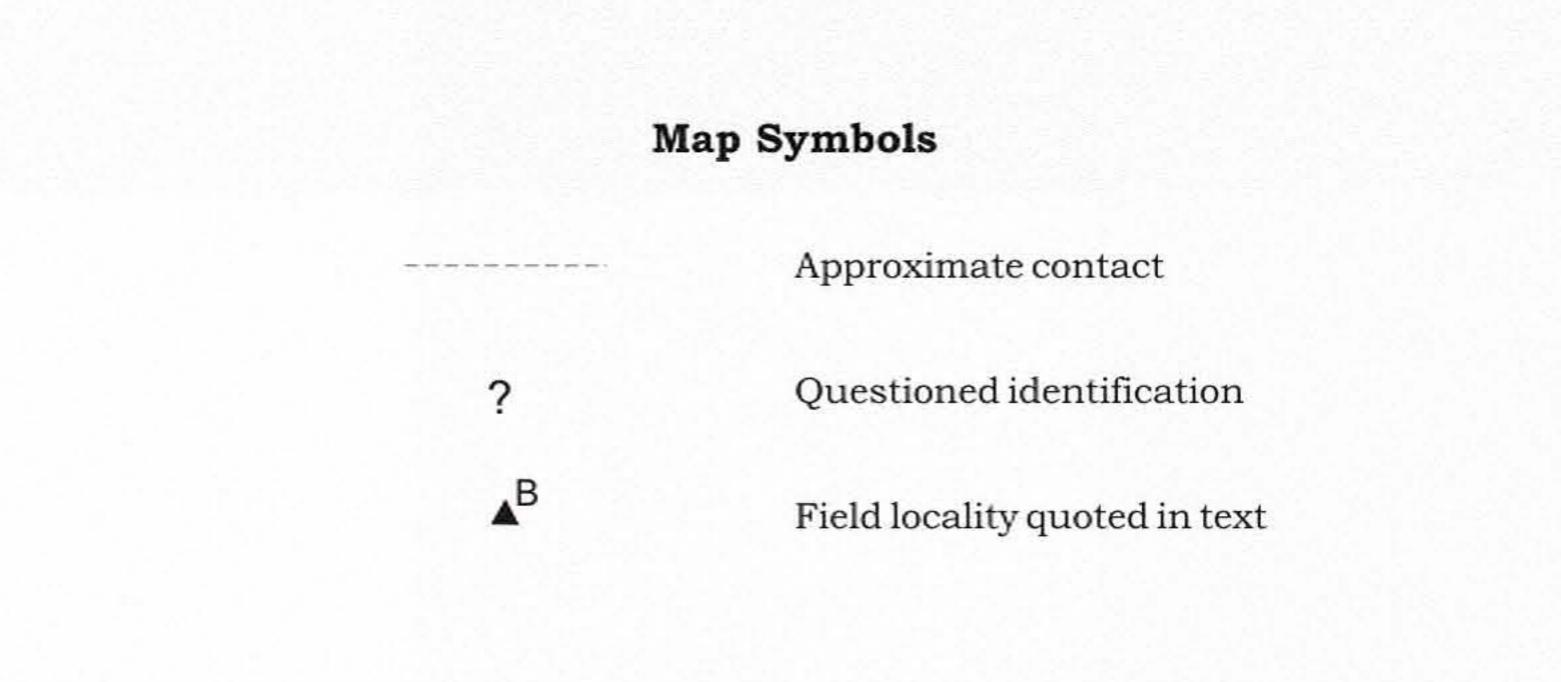
DEPOSITS OF ICE-SHOVED RAMPARTS—Chiefly well-sorted and complexly deformed polymictic pebble-cobble gravel with some medium to coarse sand pushed into a system of double 3- to 5-ft-high ridges around blown sand into ventifacts, many of which bear caliche rinds; discontinuously frozen with low ice content

SWAMP DEPOSIT—Primarily fibrous and locally woody autochthonous peat with organic silt and sand deposited in lowland sites (Kreig and Reger, 1982); up to 8 ft thick; discontinuously to continuously frozen with moderate to high ice content

QUARTZITE AND SCHIST—Chiefly quartzite with common quartz-mica schist and lesser feldspathic quartz-ite; generally garnetiferous, locally actinolitic and gneissic (Weber and others, 1977, 1978)

AUGEN GNEISS AND BIOTITE GNEISS—Chiefly medium- to coarse-grained, foliated and typically mylonitic augen gneiss with subordinate augen schist and biotite gneiss of amphibolite metamorphic facies in bedrock uplands of southern Yukon-Tanana terrane (Weber, 1971; Weber and others, 1977, 1978; Foster and others, 1973; Aleinikoff and others, 1986); generally unfrozen on south-facing slopes, sporadically to discontinuously

QUARTZ-BIOTITE GNEISS-Medium-grained, well-foliated and banded to massive gneiss interbedded with quartzite in bedrock uplands of the southern Yukon-Tanana terrane (Weber, 1971; Weber and others, 1977; 1978; Foster and others, 1973; Aleinikoff and others, 1986); weathered as deep as 50 ft; generally unfrozen on south-facing slopes, sporadically to discontinuously frozen on lower east- and west-facing slopes, and continuously frozen on north-facing slopes, all with low ice content



≈1:63,360-scale, false-color infrared aerial photographs taken in

July 1978, August 1980, and August 1981.

Electronic cartography: Alfred G. Sturmann

# Discussion

military facility of Fort Greely is being seriously considered for expansion as part of the strategic missile-

The climate in the Big Delta A-4 Quadrangle is continental, characterized by extreme temperatures ranging from the high 80s to low 90s°F in summer to as cold as -65°F in winter (Péwé and Holmes, 1964). Mean annual temperature at Delta Junction is about 27°F (de Percin and others in Péwé and Holmes, 1964). Permafrost is generally discontinuous. However, coarse-grained outwash alluvium away from modern streams isolated masses of relict permafrost (Péwé, 1955b). In near-surface silty sand capping the outwash fan north of Delta Junction, where microclimates are exceptionally cold, small ice wedges are actively forming (Péwé and Reger, 1983a). In the northern third of the quadrangle, there are continuous, ice-rich masses in retransported organic silt and sand and in peat-rich bog deposits covering treads of higher terraces (Kreig and Reger,

Annual precipitation is ≈11 inches, two-thirds of which falls as rain, and average annual snowfall is ≈3 ft. Because of its location at the junction of two large topographic corridors, the Big Delta A-4 area is subject to considerable strong-wind activity in contrast to most of the Interior (Mitchell, 1956). During a 20-yr period between 1949 and 1969 and considering only winds blowing more than about 11 mph, about 70 percent of these strong winds come from the south down the Delta River corridor during the summer in contrast to 40 to >95 percent from the east-southeast in the winter (Wendler and others, 1980). Wind velocities in excess of 95 mph have been recorded at the FAA station just south of Delta Junction (Péwé and Holmes, 1964).

In the past, strong surface winds passing across the unvegetated fluvial bars and floodplain surfaces of the Delta River picked up and spread considerable sand and silt across leeward surfaces east and west of the Delta River (Péwé, 1951; fig. 1). Eolian deposits are much thicker and widespread west of the Delta River along the east side of the Tanana River before 8,000 14C yr ago (Péwé and Reger, 1983a)(table 1, A) and in late Wisconsin time blanketed upland and lowland sites near the Tanana River with eolian sand and loess (Holmes, 1996; Holmes and others, 1996). Loess deposition continues today in the area (fig. 1). We speculate that a thin, white tephra near the base of thick loess sections along the eastern floodplain margin of the Delta River near its junction with the Tanana River is part of bed G of the Hayes tephra set H (Riehle and others, 1990; Riehle, 1994, fig. 1) [=Jarvis Ash Bed of Reger and others (1964) and Péwé (1975b) and Jarvis Creek Ash of Begét and others (1991)]. Although the Hayes tephra set has been radiocarbon dated at between 3,500 and 3,800 <sup>14</sup>C yr B.P. (Riehle and others, 1990), a series of dates in the central Alaska Range (Begét and others, 1991) places deposition of lobe G at close to 3,660 ± 125 <sup>14</sup>C yr B.P.

Major streams crossing the Big Delta A-4 Quadrangle are the Tanana River and the Delta River, which are large, glacier-fed, braided streams draining the eastern and central Alaska Range, respectively (Anderson, 1970; Dingman and others, 1971; Nelson, 1995). Active bars and the floodplain of the Tanana River are sur-Tanana River crossing (Péwé and Reger, 1983a). Based on the records taken from a single gage maintained by the U.S. Geological Survey from 1948 to 1957, the flow of the Tanana River at the highway and TAPS crossing ranged from 3,720 cfs to 62,800 cfs and averaged 14,950 cfs (Wilcox, 1980). The more intensively braided bars of the Delta River are composed of ≈60 percent gravel and ≈40 percent sand, although local surfaces are underlain by ≈90 percent sand and ≈10 percent silt (Péwé and Holmes, 1964). Periodic summer measurements of the discharge of the Delta River indicate considerable flow variation, from 24 to 9,930 cfs (Wilcox, 1980), and Dingman and others (1971) thought there was no surface flow for half the year. Jarvis Creek, a small, glacier-fed tributary stream, enters the Delta River just south of the Big Delta A-4 Quadrangle; it has a history of flooding in response to the buildup of winter icings on the floodplain. Flood waters formerly coursed down the obvious abandoned channel to enter the Delta River about 1 mi north of Delta Junction and each spring floodwaters from ice-blocked Jarvis Creek flow along the base of the high escarpment 2 mi northeast of Delta Junction (Salcha-Big Delta Soil and Water Conservation District, 1985). Flooding in 1976 left mud rinds on tree trunks up to 3 ft above ground level along the base of the escarpment.

supplies of groundwater (Péwé, 1955b; Wilcox, 1980). The regional water table slopes gently northward toward the Tanana River beneath scattered permafrost bodies and is probably recharged by water from sevimportant and dependable spawning and overwintering habitat for fish. Extensive areas of the floodplain of the Delta River are inundated by stream icings due to groundwater seepage each winter (Sloan and others,

Surface evidence of at least two major glaciations is preserved in the Big Delta A-4 Quadrangle. The type terminal moraine and associated outwash alluvium of the earlier Delta glaciation are breached by the Delta River. Till of this advance contains numerous rock types cropping out in the Alaska Range to the south. Cobbles of Alaska Range lithologies litter the surface of proximal outwash of Delta age beneath discontinuous, thin dunes of medium to coarse sand along Jack Warren Road (Péwé and Reger, 1983b), where they are pitted, faceted, grooved, and polished into classic ventifacts. A prominent linear scarp along the southern shore of Clearwater Lake between outwash of Delta age and modern floodplain deposits has been identified as a possible fault scarp by various workers in the past (for example, Weber, 1971; Carter and Galloway, 1978) but careful examination of extensive trench exposures and geophysical surveys across this lineament found no evidence of a fault (Alyeska Pipeline Service Company, unpublished data). Among others, Lindholm and others (1959, p. 40) concluded that the Clearwater Lake scarp was cut along the margin of the braided

wash fan and terrace system bordering the Delta and Tanana rivers. These glaciofluvial deposits can be traced southward up the Delta River to the type terminal moraine of the Donnelly glaciation in the Mount Hayes D-4 Quadrangle (Péwé and Holmes, 1964). A broad fan of outwash of Donnelly age and postglacial meandering streams thread across this fan surface and pass beside and between two bands of postglacial sand dunes east of the Richardson Highway (Lindholm and others, 1959). To the northwest, lateral tracing indicates that coarse outwash alluvium, which is overlain by up to 14 ft of frozen, ice-rich silt and peat to form the extensive Shaw Creek Flats (Kreig and Reger, 1982), is Donnelly in age, not Delta in age as earlier swept by strong winds into the distinctive and episodically reactivated Rosa Creek dune field north of the Big Delta A-4 Quadrangle during the Donnelly glaciation (Kreig and Reger, 1982). During this time, the outwash alluvium beneath modern Quartz Lake was exposed, and the ventifacted cobbles now found in the ice-shoved ramparts bordering the lake were fashioned by windblown sand.

The age of the Delta glaciation has long been debated. Péwé and others (1953) initially assigned the Delta

Primary evidence in the Big Delta A-4 Quadrangle for the younger Donnelly glaciation is the extensive out-

advance to the early Wisconsin glaciation, but later, on the basis of semiquantitative relative-age criteria, the Delta glaciation and its correlatives were reassigned to the earlier Illinoian glaciation (Péwé and Reger, 1983b, fig. 33). However, others working in valleys along the northern flank of the central Alaska Range and in the nearby Yukon-Tanana Upland have continued to advocate an early Wisconsin age for the Delta glaciation resolve this controversy. Reworked angular fragments of tephra at the base of unit 5 in the Canyon Creek alluvium assigned to the upper terrace of the Tanana River, which was correlated with the Delta glaciation (Weber and others, 1981, p. 177). The tephra in unit 5 of the Canyon Creek cut was later correlated with the Sheep Creek tephra in the Fairbanks area, not the Dome Ash Bed (Hamilton and Biscoff, 1984). A question nanswered by the evidence in the Canyon Creek section is how soon after its initial deposition was the Sheep Creek tephra retransported and redeposited in unit 5. Clearly, reworking occurred prior to deposition of unit 6, which is crosscut by ice-wedge casts and may represent cold conditions during the Donnelly glaciation (Weber and others, 1981, p. 177). Near Fairbanks the Sheep Creek tephra is found stratigraphica the upper Gold Hill Loess beneath the Eva Forest Bed of the last interglaciation and beneath the widespread Old Crow tephra, which is dated by the fission-track method at 140,000 ± 10,000 yr B.P. (Péwé and others, nescence method at 190,000 ± 20,000 yr B.P. (Berger and others, 1996), also confirming its deposition prior overlying outwash attributed to the Reid glaciation and, as a result, correlated the Reid glaciation with the Delta glaciation. Therefore, both the Reid glaciation in the western Yukon Territory and the Delta glaciation evidence in the central Yukon that indicates the Reid glaciation is at least as old as oxygen-isotope stage 8 lower Delta River, moraines of Delta and Donnelly ages have very similar morphologies and do not appear to be weathered much differently. Also, in the type area of the Delta glaciation, there is an inner and an outer moraine of Delta age (Péwé and Holmes, 1964). We speculate that the inner Delta-age moraine may represent a recessional moraine of late Illinoian age.

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